Long-term Studies of Snow-Vegetation Interactions

A hierarchic geographic information system helps examine links between species distributions and regional patterns of greenness

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T umans place a value on snowy alpine regions that far outweighs their small area compared with Earth's other major biomes. Mountains have long attracted people as places of spiritual and emotional refuge and physical challenge. They are scenic, contain pristine areas of high biodiversity, provide deep snowpack for skiers, and are a source of water for urban areas and agricultural regions. Scientists also recognize the importance of high mountains to regional hydrological, biogeochemical, and atmospheric processes. Alpine ecosystems are thought to be particularly sensitive to climate change and are known to have responded to such changes in the 12,000 years since deglaciation (Benedict 1970). Contributing factors for this sensitivity are the ecosystems' low productivity and tight nutrient cycling and their situation at an extreme for many plant

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Relationships among vegetation, wind, snow, and temperature regimes may help predict effects of climate change

processes (Bliss 1985).

General circulation models usually have lower confidence for predictions of precipitation than for temperature, and it is currently unclear whether global warming would lead to local increases or decreases in precipitation at high altitudes (e.g., Cess et al. 1991). There may also be a problem extrapolating from predictions at low elevations because the climates of alpine areas are often only weakly coupled to that of lowlands (Barry 1990, Greenland 1989).

Research at the Niwot Long-Term Ecological Research (LTER) site in the Indian Peaks of the Colorado Front Range focuses on the consequences of changed temperature and precipitation regimes. The distribution of snow patches and windblown areas, duration of the snow-free period, and position of meltwater drainages strongly affect the patterns of alpine plant communities. Furthermore, the hydrology of alpine watersheds responds quickly to changes in the quality of precipitation because of thin acidic soils, large volumes of snowfall, and low buffering capacity of the slow-weathering, predominantly granitic bedrock (Williams et al. 1991). One of the goals of the Niwot LTER project is thus to understand how current snowpack distributions affect patterns of vegetation from species to regional scales.

Hierarchic geographic information systems

We are using a hierarchic geographic information system to assist us in these studies. Geographic information systems (GIS) are computer hardware and software systems designed to store, manipulate, and display geographically referenced data (Star and Estes 1990). GIS databases are commonly used for multivariate analyses of spatial ecological information. Additionally, numerical models have been linked to regional and global spatial databases to develop extrapolations of ecosystem processes (e.g., Burke et al. 1991, Running et al. 1989). A hierarchic GIS (HGIS) is a nested set of GIS databases at several spatial scales. Long-term ecological studies often require data collected from a wide range of spatial domains so that, for example, changes observed in species distributions can be linked to changes in regional patterns of spectral reflectance as observed with Earth-orbiting satellites.

In this article, we demonstrate how a nested hierarchy of relatively fine-scale GIS databases can be used to help understand the links between species patterns at the level of plots, land-scape patterns of plant communities, and regional patterns seen on satellite-images. The methods described here focus on spanning the spatial domains—from that of individual plants

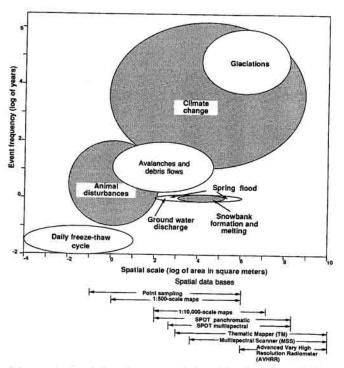


Figure 1. Spatial and temporal domains for natural disturbances at the Niwot LTER site. The shaded ellipses represent disturbance types that are major focuses of study at the site. The available data types for examining various scales of disturbance are shown at the bottom of the figure.

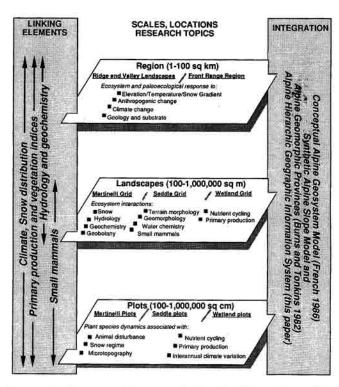


Figure 2. Conceptual framework for the Niwot hierarchical GIS.

(10⁻² m²) to that of SPOT¹ satellite images with pixel sizes of 10² m² and covering 60-kilometer-wide swaths (approximately 10⁹ m² image size). We use a suite of standard mapping methods and vegetation legends to improve compatibility between tiers of the hierarchy and promote intersite comparisons.

Conceptual framework for the alpine HGIS. Our three primary goals for establishing the alpine HGIS are to provide accurate spatial frameworks for studies of ecosystem processes and geobotanical patterns at appropriate scales, develop baselines for long-term observations of natural and anthropogenic change, and provide geographically referenced data for models of ecosystem processes. The hierarchy of disturbances in the alpine ranges from soil heave caused by the formation of needle-ice crystals at

spatial scales of 10⁻⁴–10¹ m² (Washburn 1980) to the disturbance caused by glaciers, which currently cover areas as small as 10⁴ m² but which covered areas as large as 10⁸ m² in the Front Range during the Pleistocene (Outcalt and MacPhail 1965; Figure 1). The databases are prepared to address questions related to disturbances at plot, landscape, and regional scales.

A conceptual diagram of the HGIS (Figure 2) summarizes its tiers, the major topics of research at each scale, the themes that provide linkage between tiers, and models that provide conceptual integration. At plot-level scales (10⁻²–10² m²), we are interested in the plant species dynamics associated with snow distribution. Snowpack indirectly controls the distribution of many plant species by limiting the length of the growing season (Billings and Bliss 1959, May and Webber 1982). Wind-exposed sites (Figure 3) have extremely low winter soil temperatures and high moisture stress (Ehleringer and Miller 1975, Oberbauer and Billings 1981).

The distribution of pocket gophers (Thomomys talpoides) is also strongly

controlled by snow patterns. Gophers are largely responsible for the fine-scale mosaic of many plant communities; they maintain species diversity by creating gaps in the plant canopy, redistributing nutrients and soil, and suppressing species that would otherwise dominate (Halfpenny and Southwick 1982, Osburn 1958, Thorn 1982, Willard 1979).

Our main long-term plot-level hypothesis is that plant species will react to changes in snowpack in a manner that is predictable from their present-day distribution along snow-depth gradients. In future years, we plan to test this hypothesis by artificially altering winter snow regimes with a series of large snow-fences at the Niwot LTER site. Before the experiment, we need to determine the patterns of species distribution with respect to natural snow distribution.

At the intermediate scale of landscapes (10² m²-10⁶ m²), we are examining the patterns of vegetation communities, primary production, and small-mammal distribution associated with hill-slope toposequences, and snow gradients from wind-blown sites to deep snow patches (Figure 4). Rug-

Systeme Probatoire l'Observation de la Terre-1 satellite: High-resolution visible with two modes of sensing—panchromatic blackand-white with 10-meter resolution over the range 0.51–0.73 5µm, and multispectral with 20-meter resolution in three channels (0.50–0.59, 0.61–0.68, and 0.79–0.89 5µm).

ged alpine topography and strong winds control snow distribution, resulting in highly heterogeneous, patchy landscapes (Billings 1973, 1988, Billings and Bliss 1959, Bliss 1985, Ellenberg 1988, Holway and Ward 1963, Komárková 1979, Marr 1967, May and Webber 1982, Swanson et al. 1988, Willard 1979). Abundant and complex boundaries (tension zones) are inherent in these patchy landscapes (Forman and Godron 1986), and if a climate change occurs, shifts in species composition would likely occur most quickly at boundaries and be observable at several scales (Nielson et al. 1989, Rosswall et al. 1988).

At the regional scale associated with the City of Boulder Watershed (103 km3) or the Front Range (106-108 km2), our primary focus is on the patterns of plant biomass associated with major elevation gradients (Figure 5). Also, we are interested to see if the fine-scale patterns associated with snow distribution at landscape scales have any influence on the regional trends in production. Estimates of production based on remotely sensed data provide an efficient means to examine regional patterns, yet there is relatively little research in the alpine regions, possibly due to the confounding influences of rugged topography and highly heterogeneous vegetation patterns. Our main regional hypothesis is that primary production is broadly controlled by gradients associated with changing elevation but also influenced by smaller-scale topographic interactions with wind and

Topics linking all three tiers of the hierarchy include climate, snow distribution, primary production, and spectral reflectance. The influence of small mammals is a theme that links the lower two tiers; topics related to hydrology and geochemistry link the upper two tiers. Several conceptual models and the HGIS provide a framework for the Niwot LTER project.

HGIS methods. In the Alaskan Arctic and the Colorado alpine, we are building HGIS databases that use a common suite of methods that permit comparative studies of vegetation patterns in arctic and alpine regions (Walker and Walker 1991). Most of the methods we use are described in

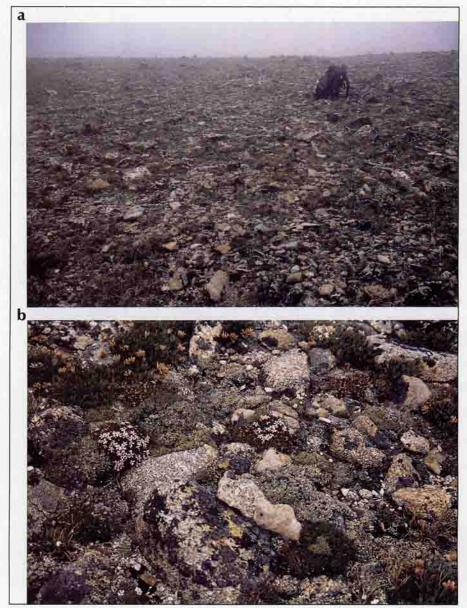


Figure 3. Fellfield vegetation. At the plot level, we are investigating the distribution of plant species in relation to snow patterns. a. Association Sileno-Paronychietum is characteristic of windblown ridges. b. The herb layer is dominated by cushion and mat-forming plants, including Paronychia pulvinata, Silene acaulis, Minuartia obtusiloba, and Trifolium nana. Photos by D. A. Walker.

the references cited in Table 1.

Geographically referencing ecological data. The LTER project is representative of many intensive ecological research programs that require topographic information and base maps at finer scales than is available from standard US Geological Survey (USGS) topographic maps. For example, the largest-scale topographic maps available for the Niwot site are at 1:24,000 scale, but the project also requires maps of the intensive research sites at 1:500 scale (Figure 6). To satisfy this requirement, an aerial-

photograph mission provided blackand-white and false-color photographs at the necessary scale.

Before obtaining the aerial photographs, it was necessary to survey ground control points to maintain horizontal and vertical accuracy across the map and to register mapped information to satellite-derived data. In most remote regions, there are few accurately surveyed benchmarks to provide control for locating spatial data. Control points can be surveyed from existing benchmarks or with the use of the global positioning system



Figure 4. Snowbeds near treeline on Niwot Ridge in the Front Range of Colorado. The relationships among terrain morphology, redistribution of snow by wind, vegetation patterns, and animal activity are the primary subjects of study at the landscape scale (10²–10⁶ m²). Photo by J. C. Halfpenny.



Figure 5. Treeline on Niwot Ridge. At the regional scale, HGIS focuses on phenomena associated with large-scale climate shifts, such as the position of the treeline and greenness of the tundra, which can be detected by means of remotely sensed data. Photo by J. C. Halfpenny.

(Michener et al. 1991, Wilkie 1990). Research grids were also surveyed for systematic field sampling and to establish a framework for georeferenced data collected at the site (Figure 7). These grids are established with reference to the universal transverse Mercator coordinate system with permanently marked grid points at 50-

meter intervals.

Orthophoto topographic maps and digital terrain models (DTMs) were constructed from the black-and-white aerial photographs taken by a professional photogrammetric engineering company. DTMs are computer files containing horizontal and vertical (x, y, z) coordinates for arrays of geo-

graphic points. DTMs are particularly useful for modeling watershed characteristics and surface energy budgets based on slope and aspect information. Three-dimensional views of the landscapes can also be made (Figure 6).

Mapping techniques. The most detailed level of HGIS is that of the individual vegetation sample plots. At the Niwot site, these plots are located at each grid point of the Niwot Ridge Saddle grid. Permanently marked vegetation plots are established at each grid point of the Saddle grid. Vegetation is monitored using a point quadrat (Figure 7). The intent of these plots is to provide a baseline for measuring long-term changes to the vegetation and microtopography and to examine species response along topoedaphic or other environmental gradients. Current plans call for repeating the sampling at five-year intervals.

GIS products provide a visually comprehensible baseline portraying the species composition at the top and bottom of the plant canopy, height of the vegetation canopy, and microtopography of the soil surface. The data allow for detailed examination of the controls of microscale environmental variation because the data are spatially referenced, both with respect to the 350 x 500-meter Saddle grid and the 1 x 1-meter permanent plots. For this study, a direct gradient analysis (Whittaker 1978) was used to examine the cover of each species along the snow gradient. The cover value for a given species was compared with the mean May snow depth for 1982, 1983, and 1984 at each of the 88 grid points. Other studies are examining the relationship of species occurrence to microtopographic variation within the plots and temporal change related to gopher disturbance.

Landscape- and regional-scale mapping. Many GIS projects use an approach in which existing mapped information is digitized into the GIS database directly from the original sources. The maps may have different map scales, map-unit resolutions, dates of data collection, and classification systems. When these different sources are combined in a GIS, artifacts may arise due to boundary mismatches and scale incompatibility (Dangermond and Harnden 1990).

Integrated geobotanical mapping

Database considerations	Method	Reference	Purpose and advantages
General considerations	(NUS) 149	(2) (2) (2) (2) (3) (3) (4) (4) (4) (4) (4) (4) (4) (4) (4) (4	W. The C. T. Annual Printer Communication of the Co
Conceptual framework	Tier diagram	Walker and Walker (1991)	Specifies questions being addressed and linkages between tiers of the hierarchy and sources of integration
Georeferencing of spatial data			
Coordinate system	Universal transverse Mercator grid (UTM)	USGS	All spatial information is referenced to a common global coordinate system Coordinates are in meters All intensive research sites within a region have the same
			geographic reference system
Research grids	Grids with 50-meter grid point spacing labeled with UTM coordinates	Walker and Walker (1991)	Provides a grid for field sampling and accurate georeferencing of all data collected at the intensive research sites Finer or coarser grids are easily constructed with reference to UTM coordinate system
Basemaps and digital terrain models (DTMs) for fine- scale databases	Orthophoto topographic maps and DTMs prepared by photogrammetry companies using high- resolution aerial photographs	This article	Provides necessary resolution for fine-scale databases
Mapping techniques			
Geobotanical mapping	Integrated mapping approach	Everett et al. (1978), Walker et al. (1981), Dangermond and Harnden (1990), Zonneveld (1988)	Based on principle that the landforms are the primary control of vegetation patterns. Vegetation map unit boundaries are guided by geomorphic features. This is particularly useful for natural landscapes with low-growing vegetation and lends itself well to an integrated GIS approach Vegetation, soils, landforms, and surface forms (geobotanical elements with the same spatial domain and often having common boundaries) are mapped simultaneously Minimizes boundary mismatches in the GIS database All elements in the database are photointerpreted at the scale of the basemap instead of combining elements mapped at different scales with varying boundary resolution
Map legends	Braun-Blanquet vegetation classification	Westhoff and van der Maarel (1976), Barkman (1986)	Names of units are based purely on floristic composition It is a hierarchic approach whereby higher level units have regional or global significance Units are formally recognized by the International Botanical Congress Structural and functional characteristics of vegetation are useful for remote sensing and modeling studies
Maps of plant species	Permanently registered 1 × 1-meter point quadrats	Mueller-Dombois and Ellenberg (1974), Walker and Walker (1991)	Detailed information regarding microtopography, structure of vegetation canopy, and species composition is provided for 100 points in each 1 × 1-meter plot Permanently registered location allows resampling of same points for long-term studies

used for many vegetation studies on Niwot Ridge (May and Webber 1982; Figure 8a). We then quantitatively determined the spatial coincidence between paired data elements (vegetation associations, snow cover, and slope) by using a frequency analysis (Evans et al. 1989; Figure 8c,d,e).

Linking ground-level observations to remotely sensed information. The methods of interfacing remote sensing and geographic information systems are now well established (Marble and Peuguet 1983, Star and Estes 1990). A 60 x 60-kilometer SPOT image of the Front Range of Colorado was obtained on 4 September 1988 (Figure 9). We used this image to analyze for trends in the normalized difference vegetation index (NDVI)2 along elevation gradients. NDVI is related to the amount of illuminated chlorophyll in the plant canopy with minimized contribution from background sources; it is often used as an index of green biomass or leaf area index (Goward et

2NDVI = (NIR - R)/(NIR + R), where NIR is the spectral reflectance in the near-infrared (0.725-1.1 mm), where light scattering from the canopy dominates, and R is the reflectance in the red chlorophyll-absorbing portion of the spectrum (0.58-0.68 5mm).

al. 1985). NDVIs have been shown to have near-linear relationships with canopy photosynthesis and conductance (Sellers 1985, 1987), and the index has been used for extrapolation of regional process estimates (e.g., Running et al. 1989, Schimel et al. 1991).

A preliminary analysis of an NDVI map of the Niwot LTER site suggested that the elevation gradient largely constrains the amount of green biomass in the alpine. This result seems intuitive because of the colder temperatures at high elevations and the prolonged snow cover. However,

	Area	-		Area
Syntaxonomic unit and author*	m² I	Percent	Nodum	Habitat
Class Elyno-Seslerietea Braun-Blanquet 1948 Order Kobresio-Caricetalia rupestris Komárková 1976 Alliance Kobresio-Caricion rupestris Komárková 1976 Association Trifolietum dasypbylli Willard 1963 Association Potentillo-Caricetum rupestris Willard 1963 Association Sileno-Paronychietum Willard 1963 Association Selaginello densae-Kobresietum myosuroidis Cox 1933 corr. Komárková 1976	7625 2 6100 36,926	4.4 0.0 3.5 21.1	Fellfield Fellfield Fellfield Dry meadow	Well-drained, basophilous to weakly acidophilous alpine areas Climax habitats on well-drained gently sloping ridge tops of the Front Range Windy, stable, cool, well-drained broad interfluves and ridges Subxeric, snow-free > 200 d Xeric to subxeric, south-facing slopes, snow-free > 200 d Xeric, extremely wind-exposed fellfields, snow-free > 200 d Subxeric to mesic turfs on gentle slopes, snow-free 150-200 d
Class Salicetea herbaceae Braun-Blanquet 1948 Order Trifolio-Deschampsietalia Komârková 1976 Alliance Deschampsio-Trifolion parryi Komárková 1976 Association Acomastylidetum rossii Willard 1963	45,659	26.1	Moist	Alpine snow patches Earlier-melting snow patches of the Front Range, snow-free 100–150 d Shallow mesic depressions and broad leeward hill slopes Mesic, early-melting snow cover
Association Deschampsio caespitosae-Trifolietum parryi Komárková 1976 Association Stellario laetae-Deschambsietum caespitosae Willard 1963	13,570	7.8	Moist meadow Moist	Subxeric to mesic, early-melting snow patches Mesic early-melting snow narches
Order Sibbaldio-Caricetalia pyrenaicae Komárková 1976 Alliance Sibbaldio-Caricion pyrenaicae Komárková 1976 Association Toninio-Sibbaldietum Willard 1963 Association Caricetum pyrenaicae Willard 1963 Association Juncetum drummondii Willard 1963 Association Phleo commutati-Caricetum nigricantis Komárková 1976 Alliance Anthelio-Pohlion obtusifoliae Komárková 1976 Association Polytrichastro alpini-Pohlietum obtusifoliae Komárková 1976	13,263 1684 1567 760 214	7.6 1.0 0.9 0.4	Snowbed Snowbed Snowbed Snowbed Snowbed	Later-melting snow patches of the Front Range, snow-free <75 d Later-melting snow patches of the low alpine Subxeric to subhygric, margins of later-melting snow Subxeric to mesic, later-melting snow patches Mesic, moderately later-melting snow patches Mesic to subhygric depressions below deep snow Bryophyte-dominated, very-later-melting snow patches Mesic to subhygric springs and later-melting snow
Class Scheuchzerio-Caricetea fuscae Tüxen 1937 Order Pediculari-Caricetalia scopulorum Komárková 1976 Alliance Pediculari-Caricion scopolorum Komárková 1976 Association Caricetum scopulorum Kiener 1939 em. Willard 1963	0688	5.1	Wet meadow	Alpine bogs and marshes Marsh communities of the Front Range, snow-free period varies Wetter, warmer, shallower snow marsh communities of the Front Range Subhygric to subhydric marshes on mineral soils
Class Betulo-Adenostyletea Braun-Blanquet et Tüxen 1943 Order Salici-Trollietalia Komårkovå 1976 Alliance Salicion planifolio-villosae Komårkovå 1976 Association Bistorto viviparae–Salicetum planifoliae Komårkovå 1976 Association Rhodiolo integrifoliae–Salicetum planifoliae Komårkovå 1976	1167	0.7	Shrub tundra Shrub tundra	Tall herb, grass, and shrub communities of the lower alpine and subalpine belt Tall herb and shrub communities of the Colorado subalpine Subxeric to subhygric willow shrublands, snow-free 100–150 d Subxeric to mesic shrublands Mesic to subhygric shrublands
Class Montio-Cardaminetea Braun-Blanquet et Tüxen 1943 Order Primulo-Cardaminetea Komárková 1976 Alliance Cardamino-Primulion parryi Komárková 1976 Association Epilobio anagallidifolii-Primuletum Komárková 1976	577	0.3	Snowbed	Montane to subalpine springs Spring communities of the Rocky Mountains, snow-free period varies Spring communities of the Colorado alpine and subalpine Subhydric to hydric springs, streams, and snow patches

*Komárková 1979. †May and Webber 1982.

0.0 13.5 0.0 100.0

56 23,610 5 174,842

Abies lasiocarpa or Picea engemannii Krummholz Barren including lichen-covered rock Water

Other map units

Total

Table 3. Optimal snow depth classes for the 50 most common species on the Niwot Ridge Saddle grid. Species are placed in the snow class where their mean cover value is the highest. Cover is calculated as percent occurrence of 8448 random points (88 plots with 96 points each; 4 points in each sample are used as registration points for the point quadrat). Frequency is the percent of occurrences in the 88 1 × 1-meter plots. Snow depth Cover class (cm) Species (percent) 0-25 Selaginella densa 2.10 23 Silene acaulis 0.77 18 Xanthoparmelia taractica 0.66 18 0.54 16 Oreoxis alpina Umbilicaria krascheninnikovii 0.38 7 0.19 3 Sporastatia testudinea 0.15 8 Arenaria fendleri 5 Paronychia pulvinata 0.14 8 Hymenoxis acaulis 0.13 Phlox sibirica 0.13 7 9 Luzula spicata 0.09 26-50 10.25 26 Kobresia myosuroides Carex rupestris 2.36 31 Cladonia pyxidata 1.59 42 1.18 41 Minuartia obtusiloba 24 Trifolium dasyphyllum 0.85 20 Lecanora rupestris 0.54 0.39 10 Rhizocarpon geographicum 11 Calamagrostis purpurascens 0.37 Thannolia subuliformis 0.25 14

Cornicularia aculeata 0.24 15 Campanula uniflora 0.15 8 Cetraria islandica 0.13 8 57 51-100 Artemisia scopulorum 3.30 Carex scopulorum 3.26 26 Salix planifolia 1.09 1 Erigeron simplex 0.84 24 Gentianoides algida 0.56 18 0.50 Salix arctica 16 Lloydia serotina 0.4120 Bistorta vivipara 0.34 11 83 101-200 Acomastylis rossii 16.08 2.46 61 Bistorta bistortoides Caltha leptosepala 1.31 16 5.55 40 201-300 Deschampsia caespitosa 27 Sibbaldia procumbens 3.27 Polytrichum piliferum 2.76 24 35 Festuca baffinensis 0.98

areas of the alpine.

Frequency (percent)

15

22

15

12

11

12

44

9

3

3

7

7

the NDVI values in 50-meter elevation classes derived from a digital terrain model. To determine if there was a systematic effect of slope or aspect, we stratified the data according to

slope-aspect classes (e.g., north, northeast, and east), and 5% slope classes. The data were also stratified according to areas east and west of the Continental Divide to assess the ef-1973).

fects of different orographic precipitation regimes. Most of the snow on the west side of the Divide occurs with storm systems from the Pacific with westerly flow, whereas the East Slope receives the largest amount of precipitation during upslope conditions with strong southerly to easterly flow (Barry Ground measurements of spectral reflectance were used to calibrate the SPOT NDVI values with mapped vegetation information in the Saddle grid

scale USGS topographic maps. NDVI

was regressed against elevation, using

was taken. Spectral reflectance was measured at each of the 88 grid points in the Saddle grid. We also obtained a clip harvest of the vegetation to determine relationship between NDVI and biomass at each grid point.

GIS. We used a field-portable radiom-

eter to obtain mean spectral reflec-

tance values for the major vegetation

types portrayed on the vegetation map.

The measurements were taken during

the same week that the SPOT image

Vegetation-snow relationships in the Colorado alpine Plant species-snow relationships. Eighty-four plant species were re-

corded at the Saddle grid plots. We examined the distribution of the 50 most common species along the snow gradient (Table 3). These species represented 77% of the cover in the 88 plots. Most of the remaining 23% was

rock (11.4%), soil (4.3%), and litter (2.5%).Many of these species are found in narrow ranges of snow depth. For example, Paronychia pulvinata is a cushion plant that has its optimal distribution on stable, dry, windblown,

rocky sites with less than 25 cm of

maximal May snow cover; it is rare in

areas with more than 50 cm of snow

(Figure 3, Figure 10a). Other plants

with a similar distribution include

Oreoxis alpina, Phlox sibirica, and

To quantify the relationships between NDVI and elevation, we stratified the alpine area above the tree line in the Front Range using the forest boundary as depicted on 1:24,000-

0.71

0.66

0.60

0.39

0.33

0.20

5.50

0.95

0.63

0.47

0.34

0.24

closer inspection of the NDVI image indicated that some large high-elevation areas are greener than areas at lower elevation, suggesting that exposure to prevailing winds and site moisture may control production over large

301-400

401-500

Lecidea atrobrunea

Castilleja occidentalis

Ranunculus adoneus

Chionophila jamesii

Lewisia pygmaea

Trifolium parryi

Carex pyrenaica

Juncus drummondii

Erigeron melanocephalus

Primula parryi

Poa arctica

Lecanora novo-mexicana

Selaginella densa. Kobresia myosuroides, although one of the most abundant plants in the Saddle (10.2% cover), has a relatively narrow distribution with respect to snow depth. Its optimal cover occurs in dry, windy sedge meadows with shallow snow cover 26–50 cm deep (Figure 10b). A large group of dry meadow species have a similar distribution pattern, including Campanula uniflora, Carex rupestris, and Trifolium dasyphyllum.

Other species are limited to sites with deep snow cover. Included here are Sibbaldia procumbens (Figure 10d), Carex pyrenaica, Chionophila jamesii, Erigeron melanocephalus, Juncus drummondii, Poa arctica, and Ranunculus adoneus. Some species, such as C. pyrenaica (Figure 10f) and E. melanocephalus, occur only in the deepest snow areas, whereas others, such as S. procumbens and Trifolium parryi (Figure 10d,e), occur across a broader range of snow cover, but they have their optimal distribution in remotely deep snow beds.

A few species span a large portion of the snow-depth gradient. The most cosmopolitan species, Acomastylis rossii, occurs in all snow depth classes, with its optimal occurrence in 1-2 m of snow (Figure 10c). A. rossii is also the most abundant species on Niwot Ridge, covering approximately 16% of the Saddle grid. Other species with wide distributions along the snow gradient (and high frequency values in Table 3) include Artemisia scopulorum, Bistorta bistortoides, Carex scopulorum, Cladonia pyxidata, Deschampsia caespitosa, Erigeron simplex, Festuca baffinensis, Minuartia obtusiloba, Polytrichum piliferum, and T. parryi. Snow plays a less important role in the distribution of these species. Other factors, such as site moisture and disturbance by gophers, may be more important to these species; nonetheless, most of these widely distributed species have a range of snow depths where their growth is optimal.

Landscape-level patterns of plant communities. The control of snow on the distribution of the dominant alpine species is also reflected at the community level. Many of the species with the clearest distributions with respect to the snow gradient are diagnostic species for the vegetation classifica-

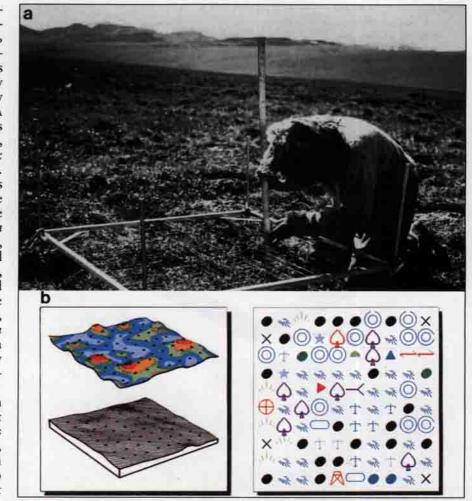


Figure 7a. Point quadrat. The metal frame has 100 gridpoints where plant species are recorded. The point quadrat can be repositioned exactly by the use of metal registration points that are nailed into the tundra. Each plot is sampled at 10-centimeter intervals. Two parallel grids of monofilament fishing line spaced 2 centimeters apart reduce the problem of parallax when sighting down on the vegetation. The quadrat frame has four small levels on each side of the frame. At each point, the observer records the species at the top and bottom of the vegetation canopy and the vertical distance of the plants from the plane of the frame. b. Example of microsite GIS database. The terrain model on the left portrays the microtopography at the base of the plant canopy (lower layer) and the height of the vegetation canopy (upper layer). The colors represent 5-centimeter height classes. The diagram on the right portrays the species composition at 10-centimeter grid point spacing at the base of the plant canopy. Separate diagrams (not shown) are made for the species at the top and bottom of the plant canopy for each grid point. The colors denote growth forms, and the shapes of the symbols denote species.

tion units, or syntaxa,³ on the vegetation map of the Saddle grid (Figure 8). For example, among the species shown in Figure 10, *Paronychia pulvinata* and *Kobresia myosuroides* (Figure 10a,b) are diagnostic taxa⁴ for alliance *Kobresio-Caricion rupestris* in

³The hierarchical units of the Braun-Blanquet vegetation classification system (class, order, alliance, and association) are collectively referred to as syntaxa in much the same way that units of species classification (class, order, family, genus, and species) are collectively called taxa.

class Elyno-Seslerietea (Komárková 1979; Table 2). This alliance includes associations that are typical of broad, well-drained, stable, windswept ridges of the Front Range. P. pulvinata has its optimal distribution in the fellfield association Sileno-Paronychietum (as described in the association tables of

In the Braun-Blanquet classification approach, diagnostic taxa are used for characterization of syntaxa; they have relatively narrow distributions within a given syntaxon even if they are not strictly restricted to that syntaxon.

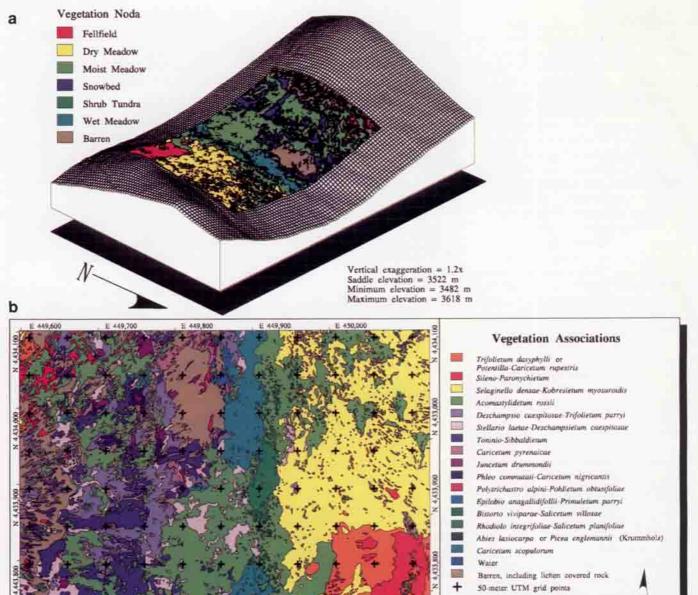


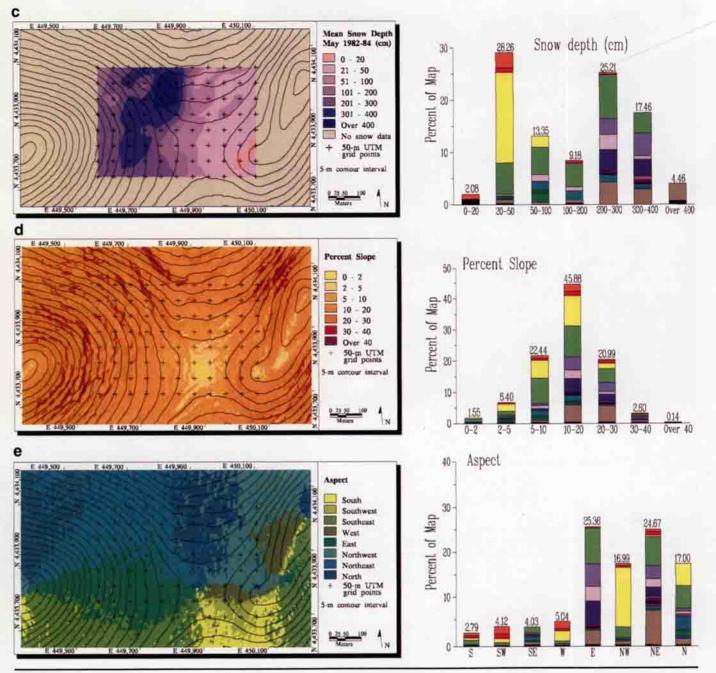
Figure 8. Vegetation, terrain, and snow depth on the Niwot Ridge Saddle grid. a. Vegetation noda (May and Webber 1982) draped over a 3-D view of the Saddle. b. Vegetation associations mapped according the classification of Komárková (1979). c. The snow map is interpolated from average May snow depths at the 88 grid points in 1982, 1983, and 1984. d and e. Slope and aspect maps derived from the digital terrain model. The bar graphs with each map portray the percentage of each map covered by the classes shown along the x-axes. Each bar is divided into sections representing the percentage of the class covered by the vegetation associations in b. The bottom cell of each bar represents the lumped total of vegetation types with area percentages too small to be shown as separate sections.

Komárková 1979). K. myosuroides has its maximal distribution in the dry meadow association Selaginello densae-Kobresietum myosuroidis. Acomastylis rossii (Figure 10c) is a widespread species that is not a diagnostic for any syntaxon, but it reaches is optimal distribution in the moist

meadow association Acomastylidetum rossii (Komárková 1979).

S. procumbens, T. parryi, and C. pyrenaica (Figure 10d,e,f) are diagnostic species for syntaxa within the snow-patch class Salicetea herbaceae. S. procumbens is a diagnostic species for the class and reaches its optimal

distribution in association Toninio-Sibbaldietum. T. parryi is a diagnostic species for the shallow-snow-patch order Trifolio-Deschampsietalia, and reaches its optimum distribution in association Deschampsio-Trifolietum parryi. C. pyrenaica is a diagnostic species for the deep-snow-patch order



Sibbaldio-Caricetalia pyrenaicae, and has its optimum distribution in the association Caricetum pyrenaicae. Numerous other diagnostic species for the syntaxa of the classes Elyno-Seslerietea and Salicetea herbaceae have distributions that are clearly related to winter snowpack and length of the growing season.

The spatial distributions of syntaxa of the classes Elyno-Seslerietea (windblown sites) and Salicetea herbaceae (snowbed sites) are strongly controlled by the distribution of snow. Syntaxa of these two classes cover 79% of the

Saddle grid (Table 2) and 78% of the alpine area on Niwot Ridge (Komárková and Webber 1978), an indication of importance of wind and snow cover in this alpine landscape. The maps of the vegetation associations and snow depths (Figure 8b and c) show a strong correspondence. Areas with less than 50 cm of snow are dominated by alliance Kobresio-Caricion rupestris. Association Sileno-Paronychietum covers 60% of the areas with less than 20 cm of snow, and association Selaginello densae-Kobresietum myosuroidis covers

69% of the areas with 21–50 cm of snow. In snow-patch areas, associations of order *Trifolio-Deschampsietalia* predominate in areas with snow depths from 50 to 200 cm (58% cover), whereas associations of order *Sibbaldio-Caricetalia pyrenaicae* and barren gravel predominate in areas with more than 200 cm of snow (45% cover).

The topographic control of these patterns is apparent (Figure 8c,d). West-facing slopes are predominantly blown free of snow. The primary vegetation types on west-facing slopes

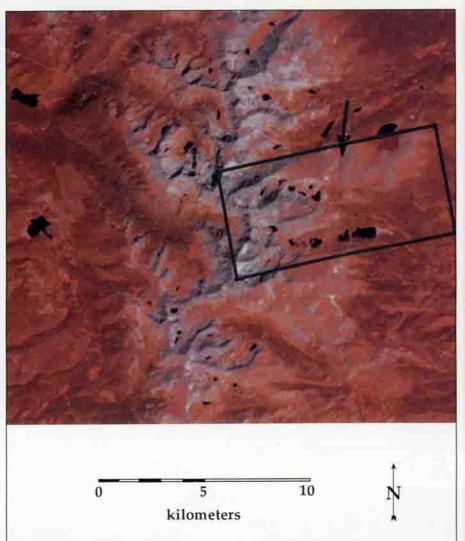


Figure 9. False-color SPOT satellite-derived image of the Front Range. The rectangle outlines the Niwot LTER site and the City of Boulder watershed. The arrow points to the Niwot Ridge Saddle area.

are either fellfield associations (Trifolietum dasyphylli and Sileno-Paronychietum) or dry sedge meadow associations (Selaginello densae-Kobresietum myosuroidis). These three associations occur on 85% of the west-facing aspects, 82% of the southwest-facing aspects, and 78% of the northwest-facing aspects. In contrast, east-facing slopes have a predominance of deep snow accumulation areas and associations of class Salicetea herbaceae (84% cover on east-facing slopes, 54% cover on southeast-facing slopes, and 58% cover on northeast-facing slopes).

Strong westerly winds are responsible for these patterns. At the D-1 weather station, 2 km west of the Saddle, 79% of the hourly mean wind directions were from the west and the

mean wind velocity was 11 m sec⁻¹ during the winter months (October-April) from January 1988 to December 1989. Maximal wind speeds exceeded 30 m sec⁻¹ during all winter months. During the same period, winds were somewhat more variable at the Saddle due to topographic steering of winds around the west knoll of the Saddle, but still more than 81% of the winds were out of either the west or northwest.⁵

The derived slope and aspect maps (Figure 8c,d) indicate considerable topographic variation at scales finer than the 50-meter resolution of the Saddle grid used to make the snow map. This topographic variation undoubtedly affects snow distribution at

a finer scale than is portrayed in Figure 8c. The correspondence between snow depth and vegetation would, therefore, likely be even stronger if the snow were mapped at the same resolution as the vegetation.

Regional patterns of NDVI. Comparison of NDVI values with snow-depth and soil-moisture values suggest an underlying relationship between soil moisture and vegetation greenness, as inferred by NDVI (Figure 11b). Within the Saddle grid, NDVI shows a trend of increasing value from barren to fellfield to dry meadow to snowbed to moist meadow to wet meadow to shrub. High NDVI values in the shrub and wet meadow areas are due to lateseason greenness resulting from moisture supplied by the late-melting snow. The low NDVI values for fellfields (associations Sileno-Paronychietum and Trifolietum dasyphylli), dry meadows (association Selaginello densae-Kobresietum myosuroidis), and windswept, primarily west-facing, areas in general (i.e., alliance Kobresio-Caricion rupestris) are likely due to low production caused by droughty site conditions. In contrast, moist meadows typical of early-melting snow patches on east facing slopes (alliance Deschampsio-Trifolion parryi) have relatively high soil moisture and high NDVI values. Late-melting snow patches (alliance Sibbaldio-Caricion pyrenaicae) are covered by snow during much of the growing season and experience a prolonged melt that washes away finer particles to leave a rocky bed with low production and low NDVI.

Regional patterns of NDVI along elevation gradients are strongly influenced by the contrast between eastand west-facing slopes. Regressions of NDVI versus elevation show the expected general decrease in NDVI with increasing elevation, with considerable scatter about the mean due to landscape heterogeneity (Figure 12a-c). This trend applies to almost all slopes and aspects on both the east and west side of the Continental Divide. The only aspects showing no trend in NDVI with elevation are westfacing aspects on the east side of the divide, probably due to the influence of strong westerly winds that constrain production on these sites at all elevations (Figures 3, 12d). The lack

⁵Niwot LTER, unpublished data, 1988-1989.

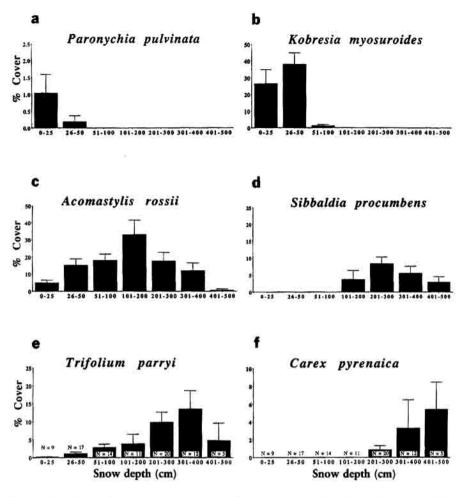


Figure 10. Distribution (plus or minus standard error) of six species along the snow gradient. Species cover values are from the 88 1 x 1-meter plots at the gridpoints of the Saddle grid. Snow measurements are the means of maximal snow depths measured at the gridpoints in May of 1982, 1983, and 1984. N = number of plots in each snow-depth class.

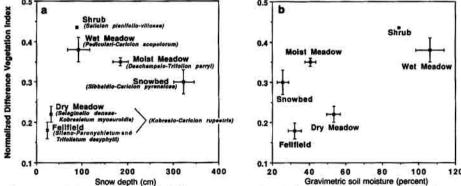


Figure 11. Mean normalized difference vegetation index for the vegetation noda on the Saddle grid (plus or minus standard errors). Noda are arranged along a snow gradient (a) and along a soil moisture gradient (b). The relatively high gravimetric soil moisture values for dry meadows compared to moist meadows is due to high organic content (low bulk density) of the Kobresietum soils.

of a similar relationship on west-facing aspects west of the Continental Divide suggests that a different wind regime exists on the West Slope. Although there are no climate stations on the west side of the Divide in the Indian Peaks, other investigators have observed that krummholz and stands of *Sileno-Paronychietum* are uncommon west of the Divide, suggesting

generally calmer conditions (Komár-ková 1979). East-facing slopes on the east side of the Divide have the steepest regression, possibly due to the generally moister soil conditions at low elevations and the constraining influence of later-melting snowpack at higher elevations. The presumably calmer conditions on the west side of the Divide result in less contrast between east- and west-facing slopes, although the y-intercept and the slope of the regression curve are greater for the east-facing slopes on both sides of the Divide.

The contrast in the regression lines between the east and west sides of the Divide suggest that the NDVI-versuselevation relationship could be used to compare the effects of different climates on regional patterns of greenness in the alpine. For example, in a humid summer climate such as the White Mountains of New Hampshire, we might find generally higher NDVI values at all elevations above the treeline, causing a shift of the regression line upward (assuming that the xaxis portrays elevation above the treeline, instead of above sea level). On the other hand, mountain ranges with dry summer climates, such as the Sierras, may have relatively low NDVI for all aspects at all elevations. A combination of warmer summer temperatures and higher precipitation may shift the line to the right. On mountain ranges with light winds, such as some areas of the Brooks Range, climates may show little contrast between windward and leeward aspects; whereas windy ranges such as the Big Horns should show a strong difference in the slopes of the regression lines on wind-influenced aspects. Similar logic could be used to predict the consequences of various climate change scenarios in the Front Range. Of course, substrate would also have a strong influence on site moisture regimes, and it would be interesting to use mountain ranges with a variety of bedrock types to examine the influence of bedrock on the NDVI-versuselevation relationship.

Linking pattern to process. The hierarchic analysis of alpine vegetation patterns emphasizes that plant species, community, and green-biomass patterns are largely controlled by snow distribution, which varies greatly with

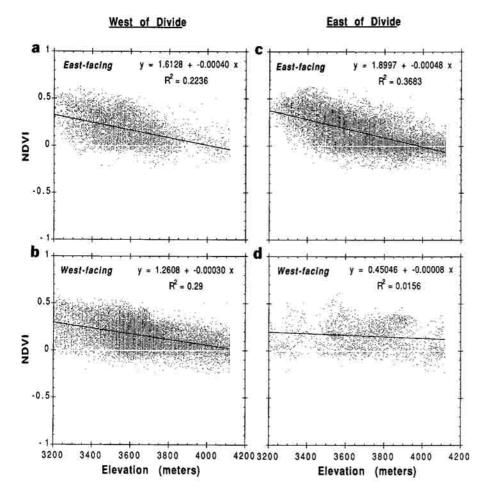


Figure 12. Normalized difference vegetation index versus elevation for two aspect classes on 6-15% slopes on the west and east sides of the Continental Divide; p <0.001 for all regressions.

topography and wind patterns. The snow regime affects water availability, soil and leaf temperature, surface microclimate, and timing and duration of the growing season. Because snow meltwater is a major source of nutrients to nitrogen- and phosphorus-deficient alpine systems, landscape-level patterns of available soil nutrients are also likely to be closely linked to snow distribution patterns (Bowman 1992, Reddy and Caine 1990). Modification of snow and moisture regimes through variations in climate or direct anthropogenic impacts may significantly alter nutrient and vegetation patterns and, as a consequence, landscape-level processes.

Specific questions currently being addressed by research at Niwot Ridge include: Will changing snow regimes influence the spatial distribution of nutrients and directly affect ecosys-

tem-level nutrient fluxes? How will alpine ecosystems adjust to altered snow regimes, and how will these adjustments manifest at different levels of ecosystem organization? The answers to these and other processlevel questions are being addressed with an experiment in which we are altering snow regimes through the use of snow fences. A prediction of the effects of future altered precipitation regimes requires a hierarchical approach to examining alpine ecosystem response, comparative studies in mountain ranges with other climates, and experimental research on the effects of altered precipitation regimes.

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